

The Ocean in near Equilibrium with Respect to Atmospheric CH₃Br

Lei Hu, Shari Yvon-Lewis, Yina Liu,
Thomas S. Bianchi

Acknowledgements:

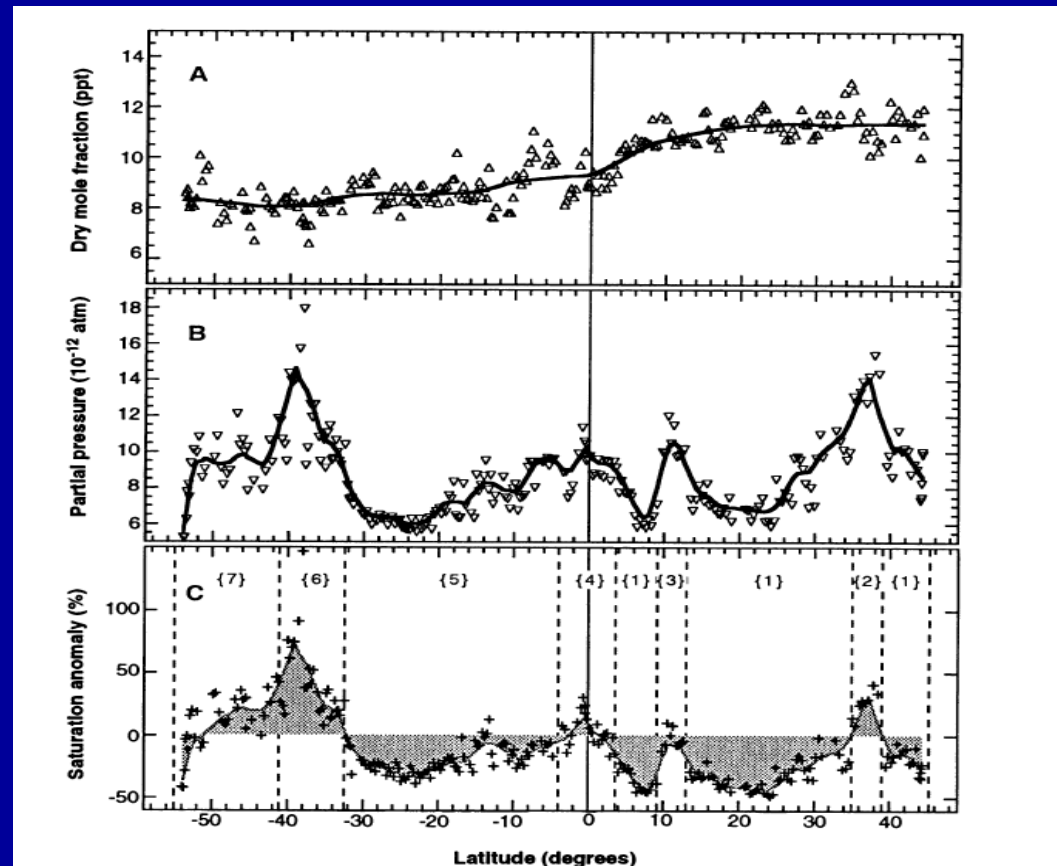
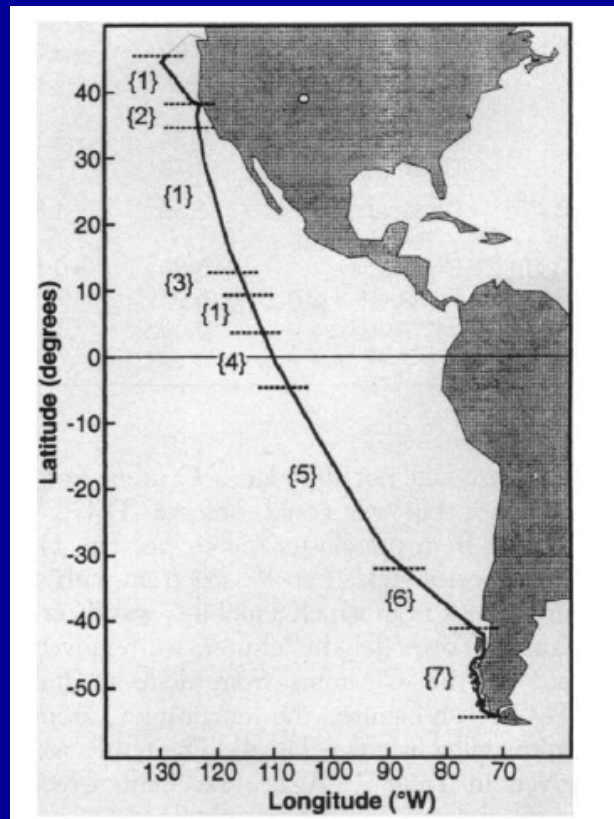
- Dr. Jim Butler and Dr. Steve Montzka
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 - OCE-0927874.
- The captains and crews of the R/V Thomas Thompson and FS Polarstern.



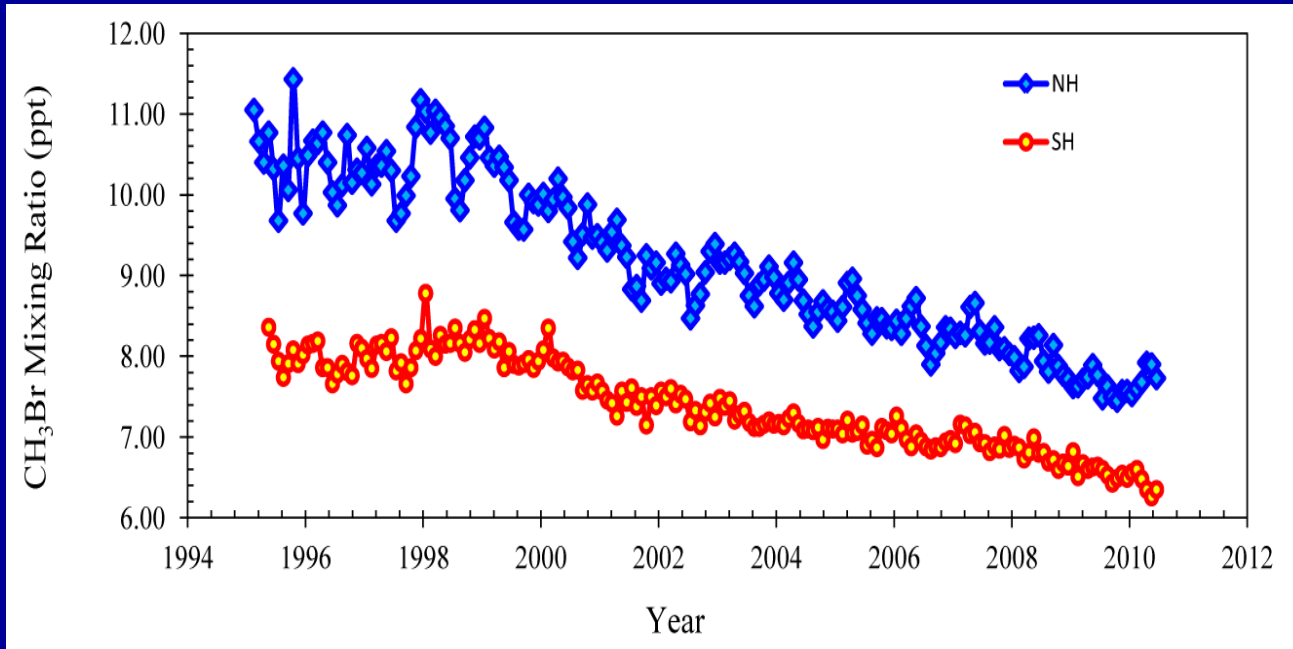
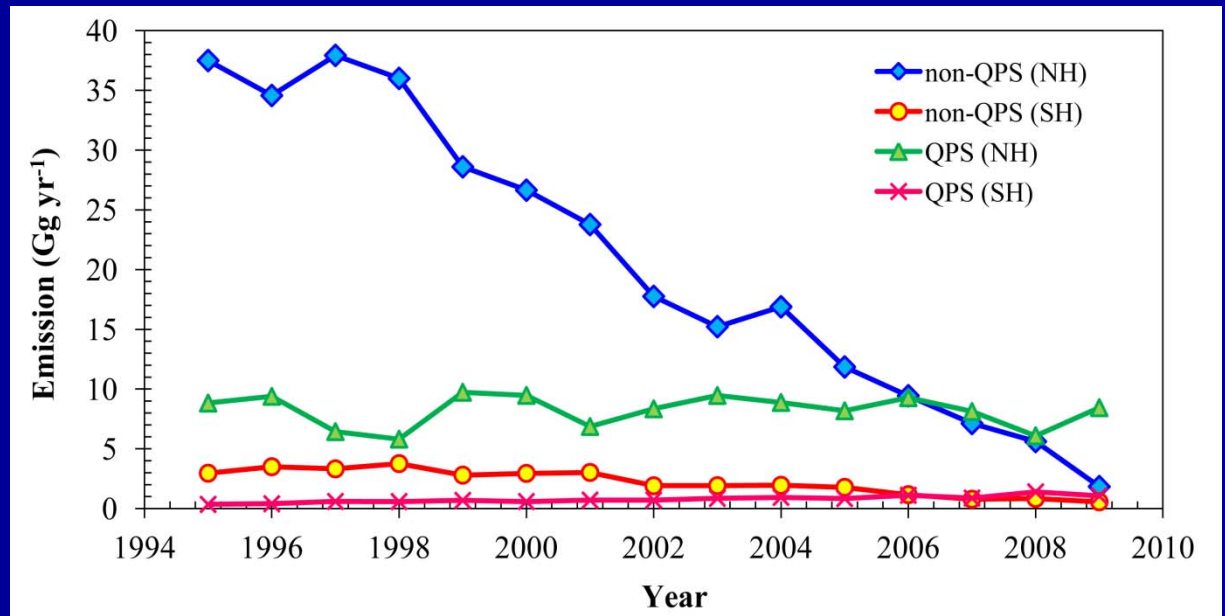
A Net Sink for Atmospheric CH_3Br in the East Pacific Ocean

Jürgen M. Lobert, James H. Butler, Stephen A. Montzka, Laurie S. Geller, Richard C. Myers, James W. Elkins

BLAST I, 1994
Published in
Science in 1995



Phase-Out



Atmospheric
Observations from
NOAA/ESRL Global
Monitoring Division

Buffering Effect of the Ocean

GEOPHYSICAL RESEARCH LETTERS, VOL. 21, NO. 3, PAGES 185-188, FEBRUARY 1, 1994

The potential role of the ocean in regulating atmospheric CH_3Br

James H. Butler

Climate Monitoring and Diagnostics Laboratory, National Oceanic and Atmospheric Administration,
Boulder, Colorado

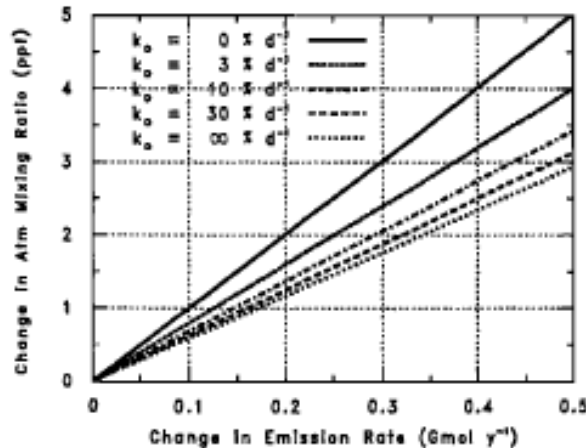
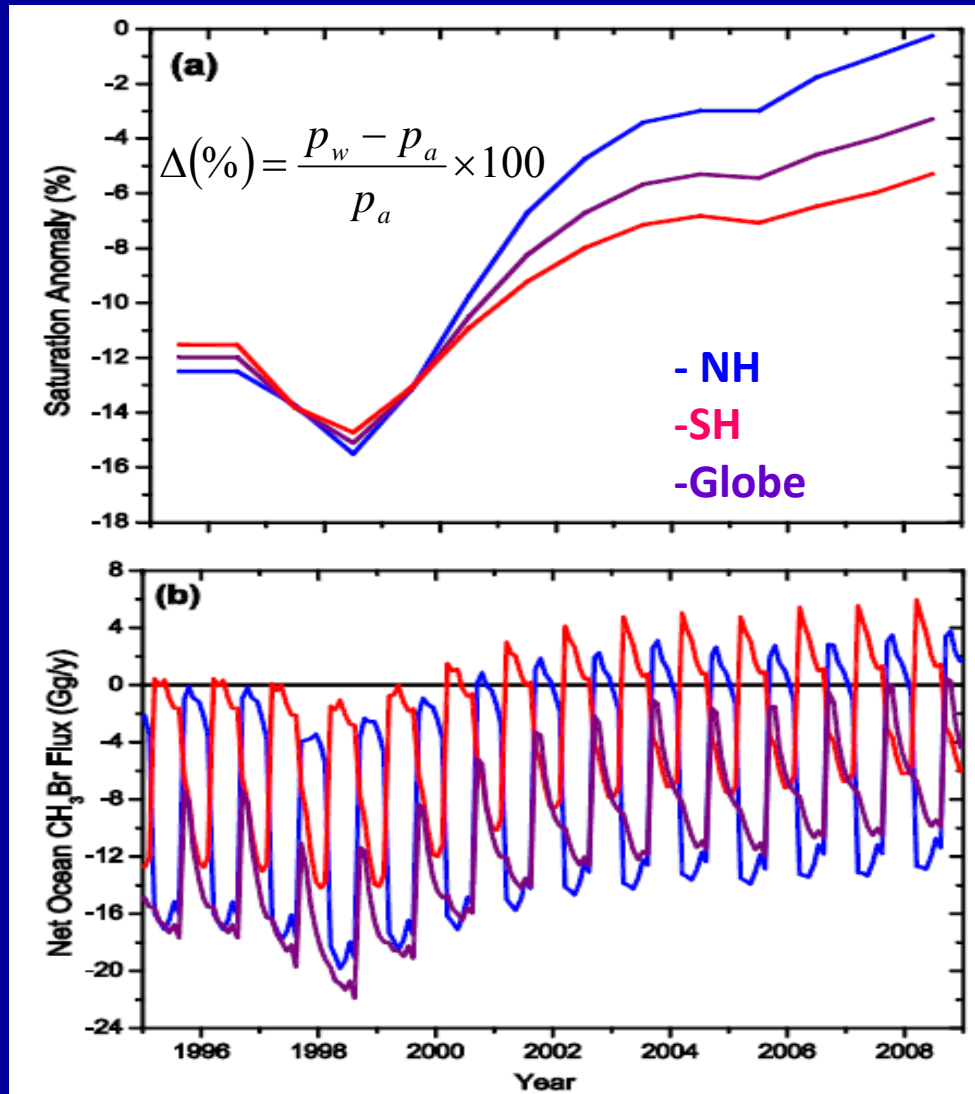


Fig. 2. Steady-state change in atmospheric mixing ratio (ppt) resulting from a changes in anthropogenic emission rate (Gmol y^{-1}). Curves are calculated from equation 12, with the current best estimates for k_a , k_1 , and k_2 as given in Table 1, and allowing k_o to vary from 0 to ∞ ($R = 0$ to 1). The curve for $k_o = 0$ represents no oceanic loss, whereas for $k_o = \infty$ the oceanic degradation rate is infinitely fast.

Ocean should become
less undersaturated as
atmospheric
concentration decreases

Recent Models Predicted Less Negative Saturation State



1996:

$$\Delta(\%) = -11.7\%$$

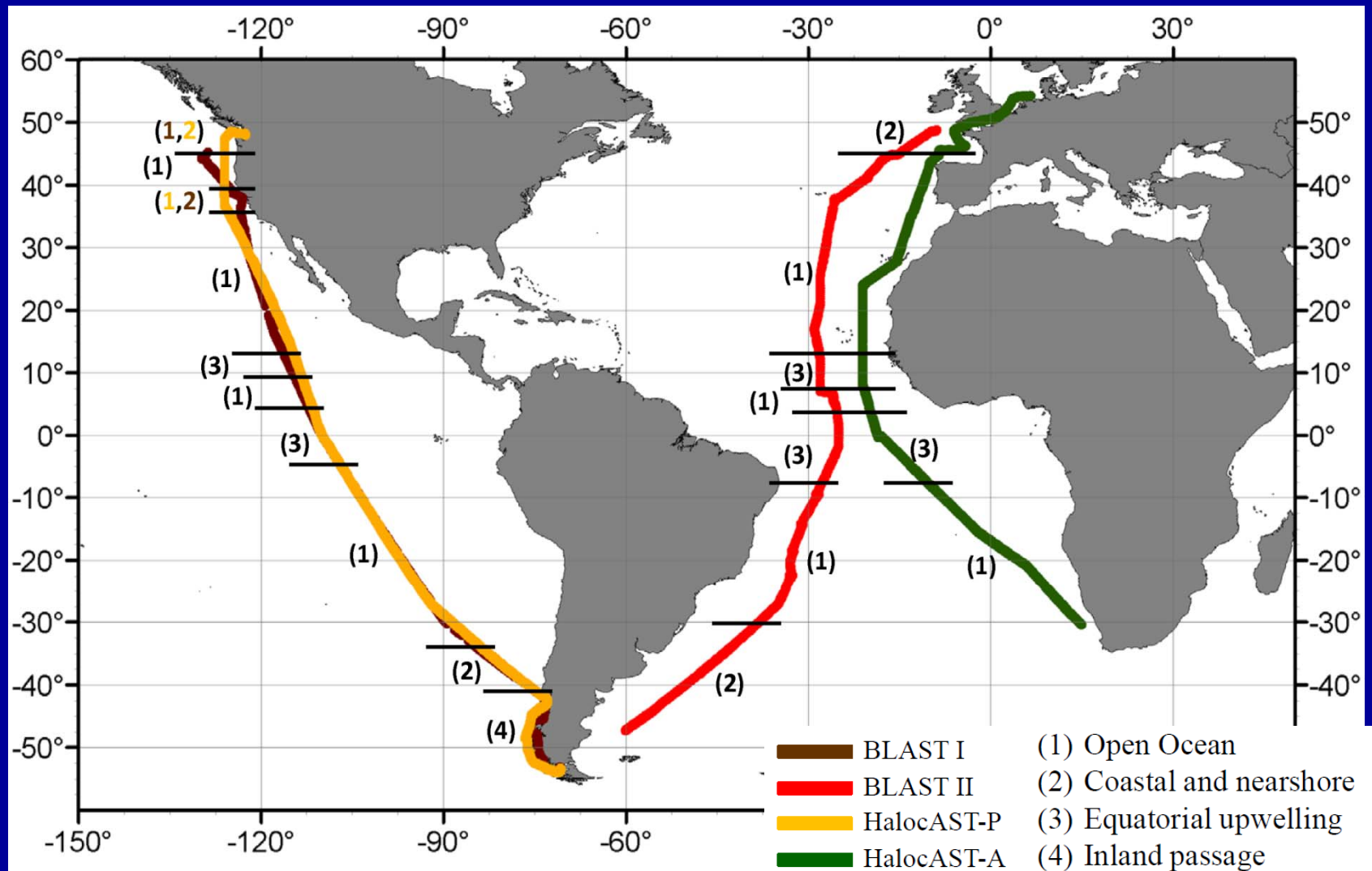
Net sea-to-air flux:
-14 Gg yr⁻¹

2007:

$$\Delta(\%) = -6.0\%$$

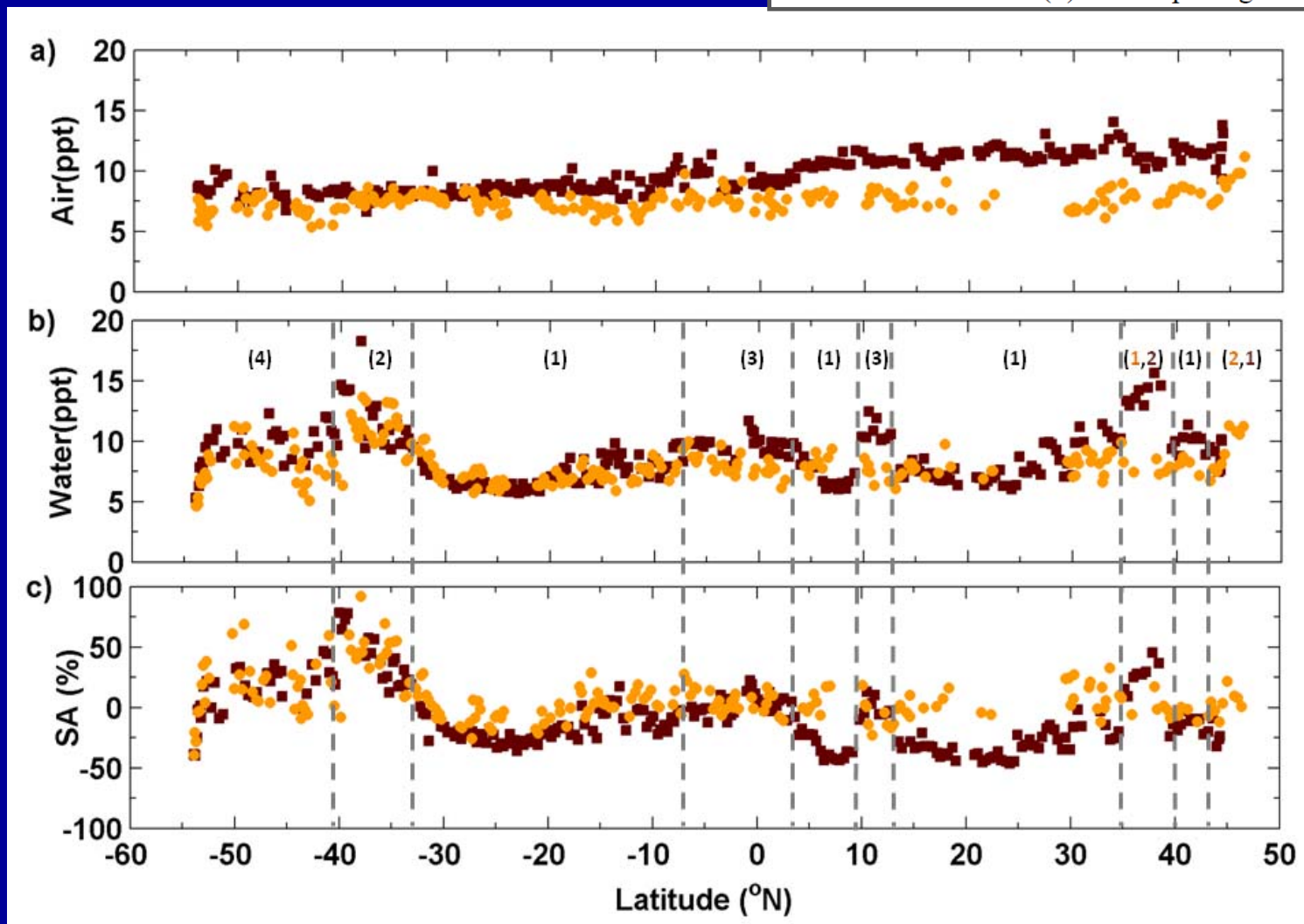
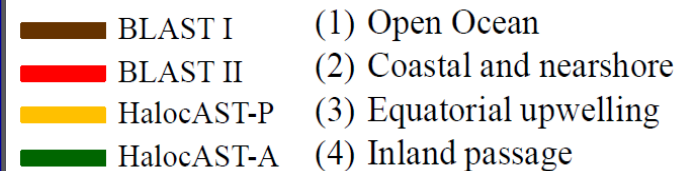
Net sea-to-air flux:
-6.6 Gg yr⁻¹

Halocarbon Air-Sea Transects – Pacific and Atlantic



Hu, et al., Submitted, GBC

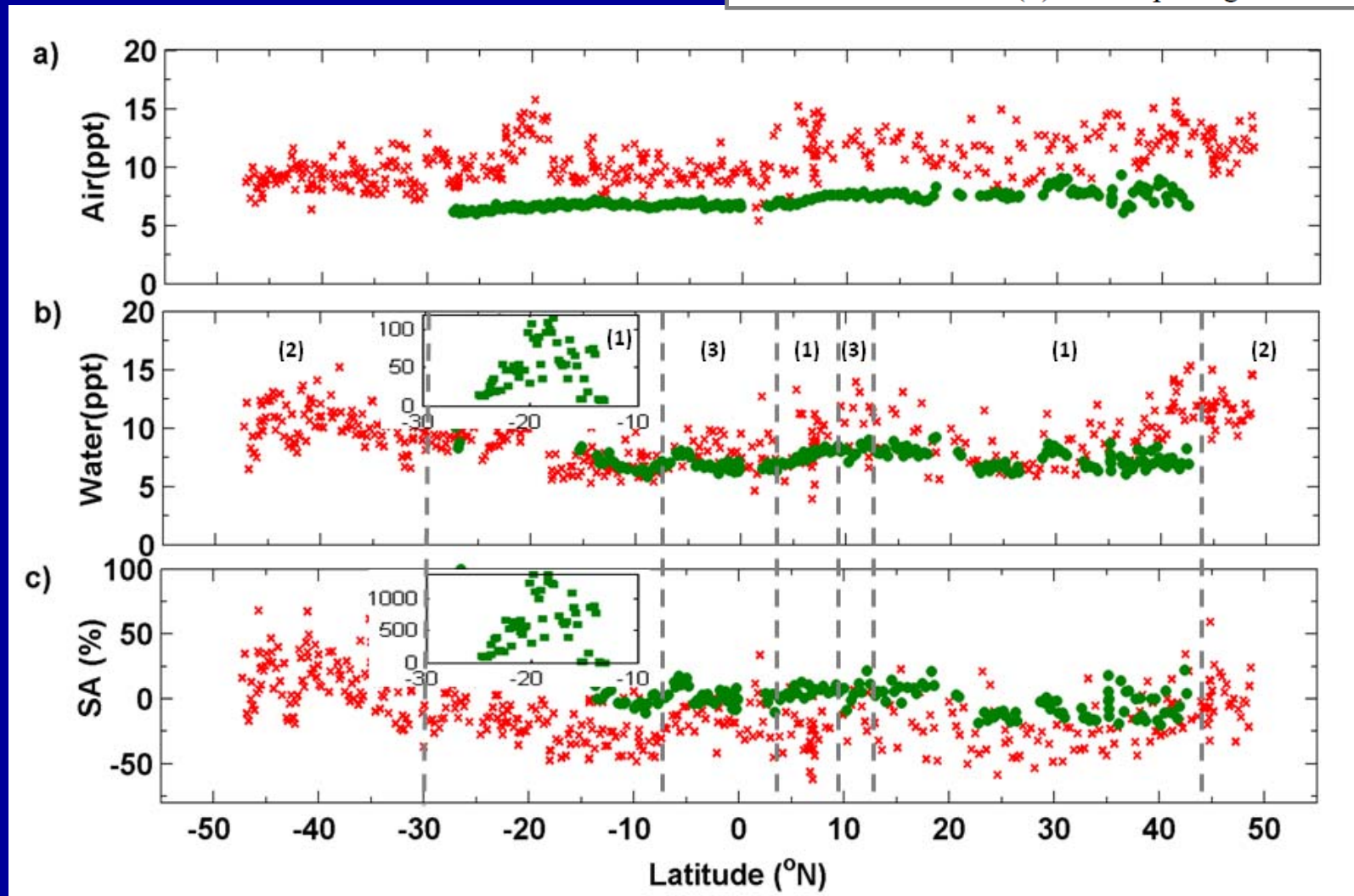
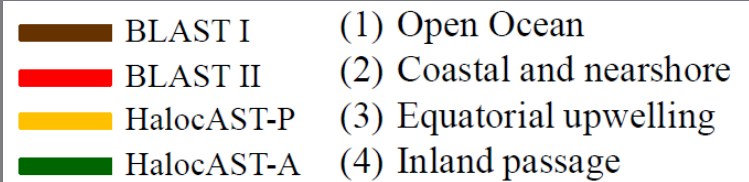
CH₃Br data from BLAST I and HalocAST-P



Hu, et al., Submitted, GBC

NOAA/ESRL GMAC 2012

CH₃Br data from BLAST II and HalocAST-A



Hu, et al., Submitted, GBC

NOAA/ESRL GMAC 2012



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Global Extrapolation

Region	Area Weighting Factor	$\Delta_{\text{CH}_3\text{Br}}$ (%)	Flux (Gg yr^{-1})	Production (Gg yr^{-1})
Open Ocean	0.8	0.3	-0.05	1.2×10^2
Coastal	0.1	29.5	2.6	15
Upwelling	0.1	2.4	0.09	21
Global (HalocAST)		3.4	2.6	1.5×10^2
Global (BLAST)		-15.7	-12.6	1.5×10^2

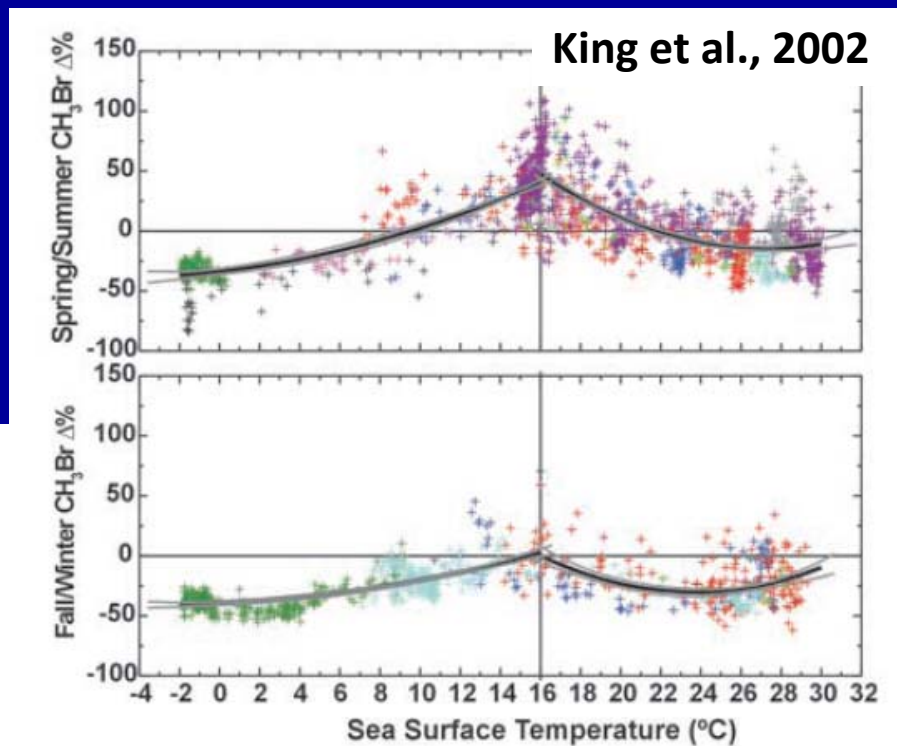
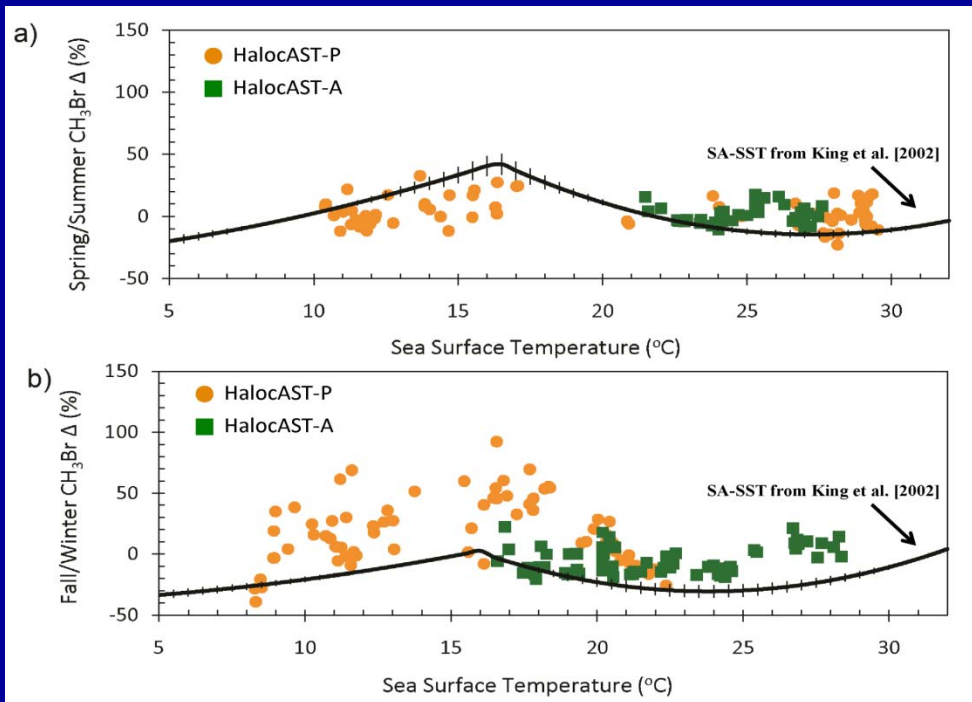
Problems with global extrapolation of fluxes:

- Extrapolated fluxes may be biased by regional saturation anomalies or regional in-situ wind speeds.

$\Delta(\%)$ - SST relationships before the phase-out is no longer valid

$\Delta(\%)$ - SST relationships before the phase-out

$$F = k_w H \frac{\Delta}{100} p_a$$



$\Delta(\%)$ - SST relationships and HalocAST data

Net Flux (F) = Emission (E) – Uptake (U)

18 year record of atmospheric CH_3Br from NOAA/ESRL GMD

$$U = f(p_a, k_{bio}, \text{windspeed}, SST, Sal, MLD)$$

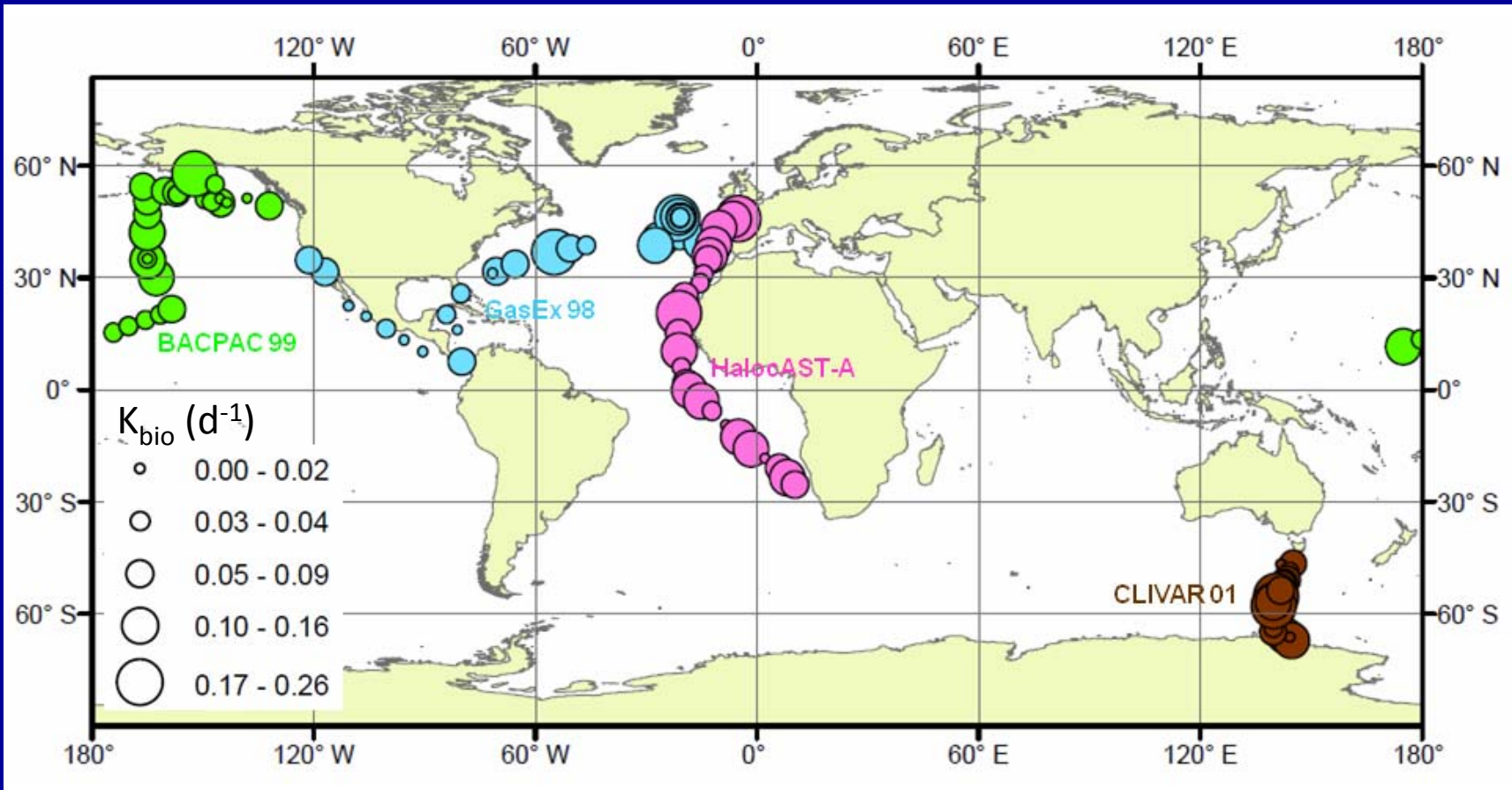
Measured before and after phase-out

DS279: $1^\circ \times 1^\circ$ gridded datasets of wind speed, SST, Sal, MLD

$$E = f(P, k_{bio}, \text{windspeed}, SST, Sal, MLD)$$

$$P = f(p_a, k_{bio}, SST, WindSpeed, Sal, MLD)$$

Biological Degradation Rate Constants



- No significant difference between degradation rate constants observed in the northeastern Atlantic during GasEx 98 and HalocAST-A.
- Global mean biological degradation rate constant = $0.05 (\pm 0.01 \text{ S.E.}) d^{-1}$.

1° x 1° Ocean Model

The old 1° x 1° Ocean Model (e.g. Yvon and Butler, 1996; King et al., 2002; Yvon-Lewis et al., 2009)

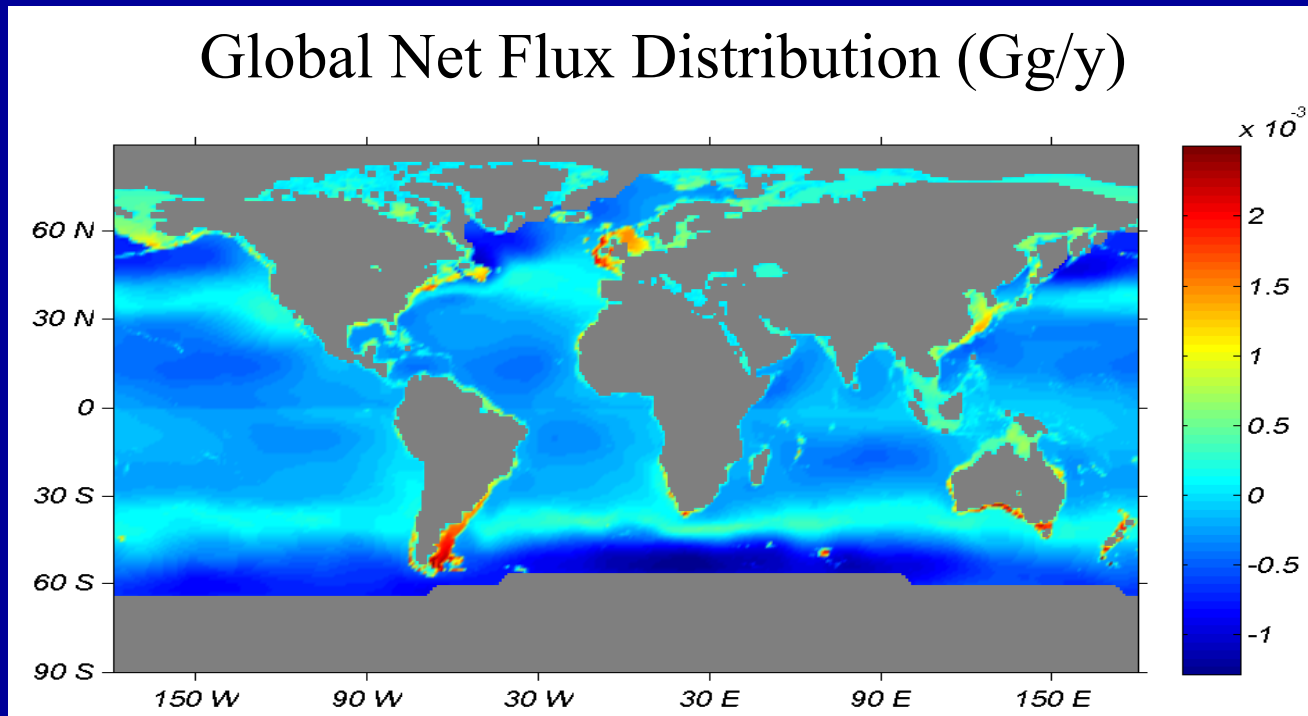
- Used Wanninkhof [1992].
- Did not include separate coastal rates and used only open ocean values for production rates, biological degradation rate constants, etc.

Revised 1° x 1° Ocean Model

- ✓ Uses Sweeney et al. [2007].
- ✓ Includes rates for both the coastal and open ocean regions.

Before the CH₃Br Phase-Out

	Emission (Gg/y)	Uptake Rate (Gg/y)	Net Flux (Gg/y)
Global Ocean (W92)	42	-56	-14
Global Ocean (S07)	31	-41	-10
Global Ocean (S07 +)	34.2	-41.1	-6.9
Coastal Ocean	4.6	-2.4	2.2
Open Ocean	29.6	-38.7	-9.1



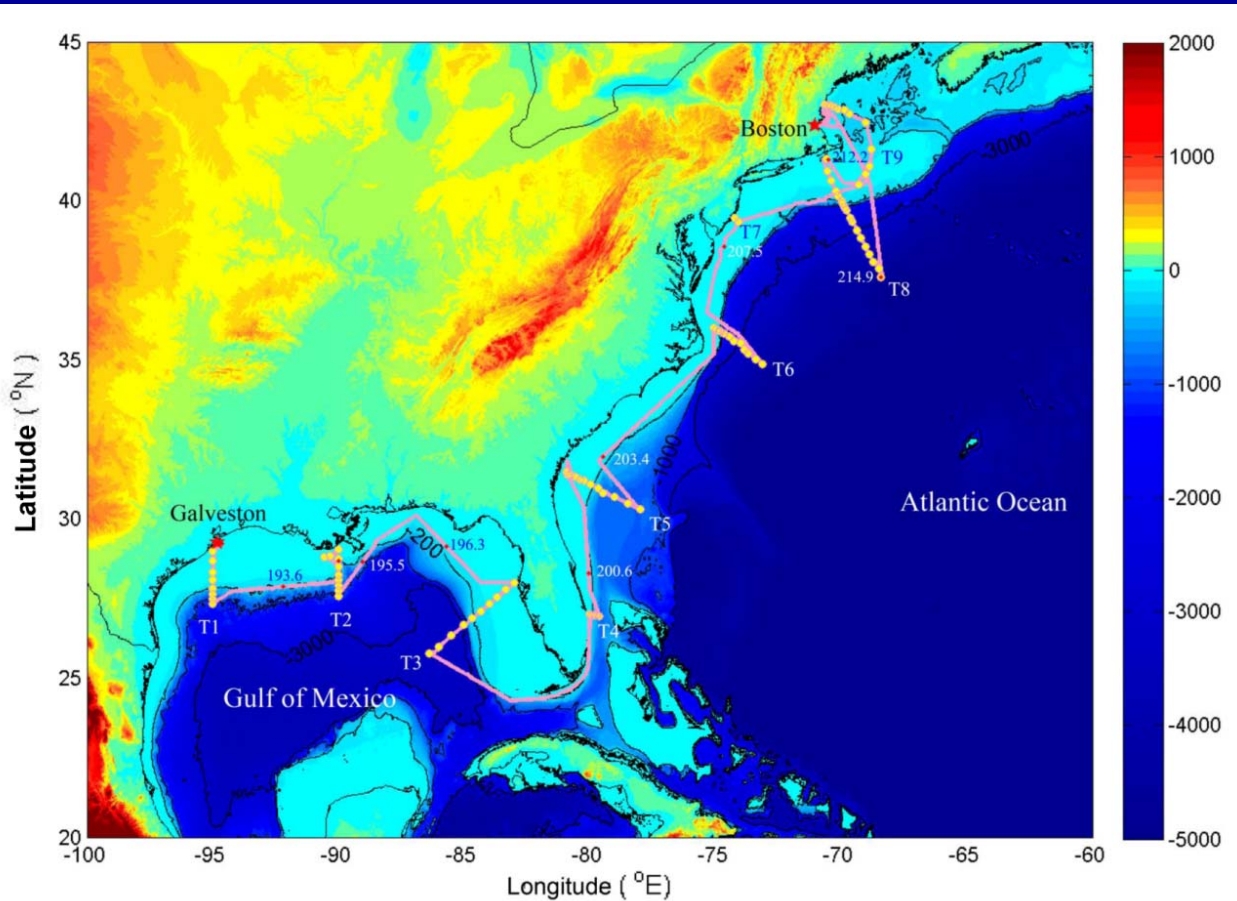
At the End of the non-QPS Phaseout (2010)

	Emission (Gg/y)	Uptake Rate (Gg/y)	Net Flux (Gg/y)
Global Ocean (Cruise Data Extrapolation)	?	?	<u>2.6</u>
Coastal Ocean	?	?	2.6
Open Ocean	?	?	0
Global Ocean (Model)	34.2	-31.5	<u>2.7</u>
Coastal Ocean	4.6	-0.8	3.8
Open Ocean	29.6	-30.7	-1.1

Summary and Conclusions

- ◆ Atmospheric methyl bromide concentrations measured during the recent cruises are consistent with the NOAA flask network data.
- ◆ The degradation rate constants measured in 2010 are consistent with previous measurements.
- ◆ The observed global mean saturation anomaly increased to $\sim 3.4\%$
- ◆ The oceans became a small net source of CH_3Br to the atmosphere with a global mean flux of $\sim 3 \text{ Gg/y}$

Differences Between Coastal-Oceanic Areas and Open-Oceanic Areas



• Production Rates:

Mean production rate in coastal areas of HalocAST and GOMECC: $0.62 \text{ nmol m}^{-3} \text{ d}^{-1}$.

Global open-ocean production rate: $0.15 \text{ nmol m}^{-3} \text{ d}^{-1}$ (Yvon-Lewis et al., 2004)

• Biological degradation rate constant

Coastal ocean (King et al., 1997): 0.09 d^{-1}

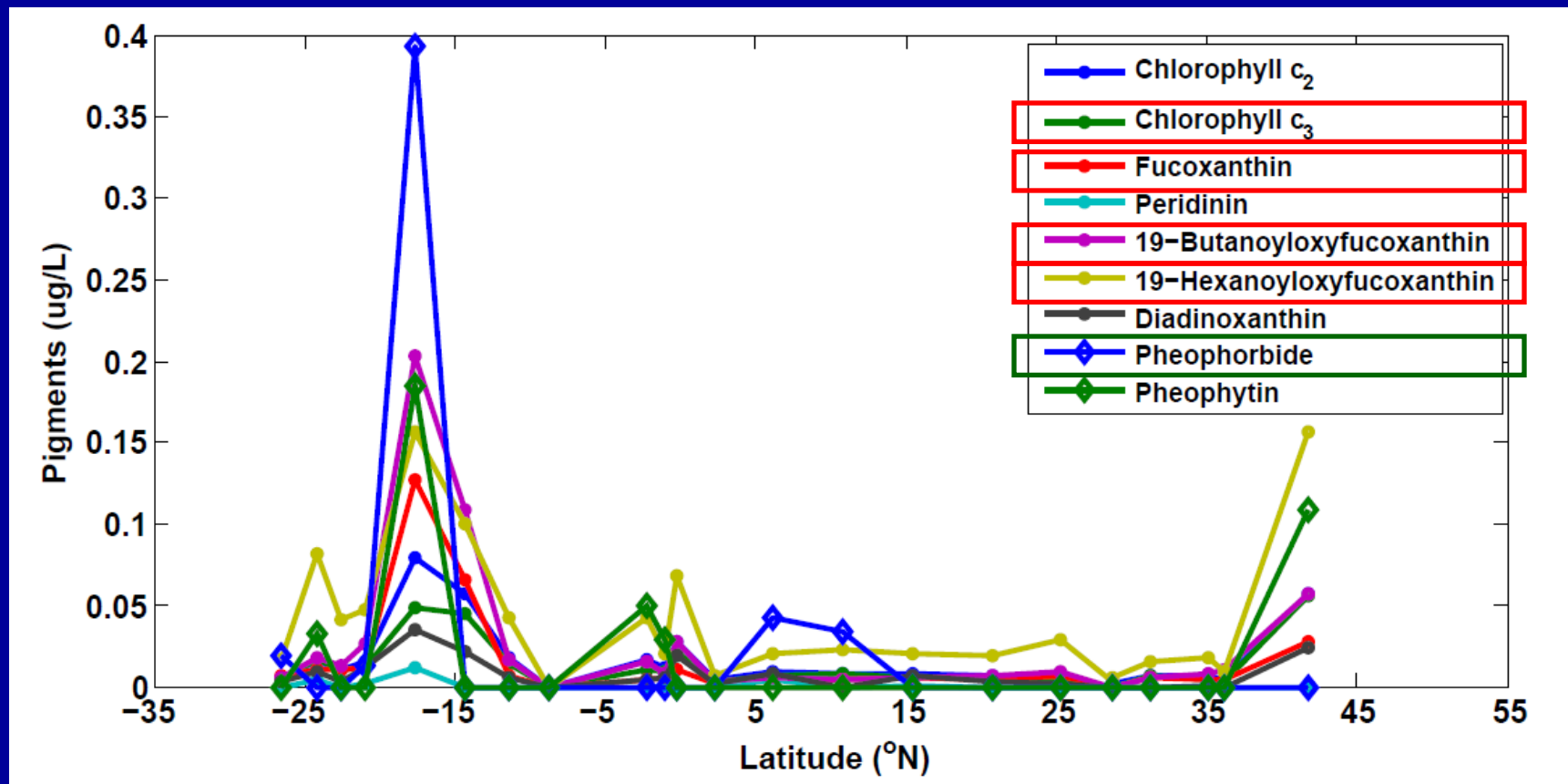
Global open ocean mean: 0.05 d^{-1}

Cruise Track of GOMECC
(Hu, et al., 2010, GBC)

Atmospheric Methyl Bromide Budget (Gg yr⁻¹)

	1996 - 1998	Range	2008	Range
SOURCES				
Fumigation - dispersive (soils)	41.5	28.1 to 55.6	6.5	4.6 to 9.0
Fumigation - quarantine/pre-shipment	7.9	7.4 to 8.5	7.6	7.1 to 8.1
Ocean	42	34 to 49	42	34 to 49
Biomass Burning	29	10 to 40	29	10 to 40
Leaded Gasoline	5.7	4.0 to 7.4	<5.7	
Temperature peatlands	0.6	-0.1 to 1.3	0.6	-0.1 to 1.3
Rice Paddies	0.7	0.1 to 1.7	0.7	0.1 to 1.7
Coastal Salt Marshes	7	0.6 to 14	7	0.6 to 14
Mangroves	1.3	1.2 - 1.3	1.3	1.2 - 1.3
Shrublands	0.2	0 to 1	0.2	0 to 1
Rapeseed	4.9	3.8 to 5.8	5.1	4.0 to 6.1
Fungus (Litter Decay)	1.7	0.5 to 5.2	1.7	0.5 to 5.2
Fungus (Leaf-cutter Ants)	0.5		0.5	
Subtotal (Sources)	143		111.5	
SINKS				
Ocean	56	49 to 64	49	45 to 52
OH and Photolysis	77		63.6	
Soils	40	23 to 56	32	19 to 44
Subtotal (Sinks)	177		147.6	
Total (SOURCES - SINKS)	-34		-36.1	

Pigments that correlated with seawater CH₃Br during 13 to 25 °S



- Elevated CH₃Br was associated with two main algal groups, prymnesiophytes and dinoflagellates.
- *Emiliana huxleyi* and *Phaeocystis sp.* can produce CH₃Br at a significant rate.
- It is likely that elevated CH₃Br was at least partly associated with *phaeocystis sp.*, which were grazed by zooplankton, or at the senescent stage or underwent autolysis.