

# Updated Aerosol Climatology for Cape Point, South Africa

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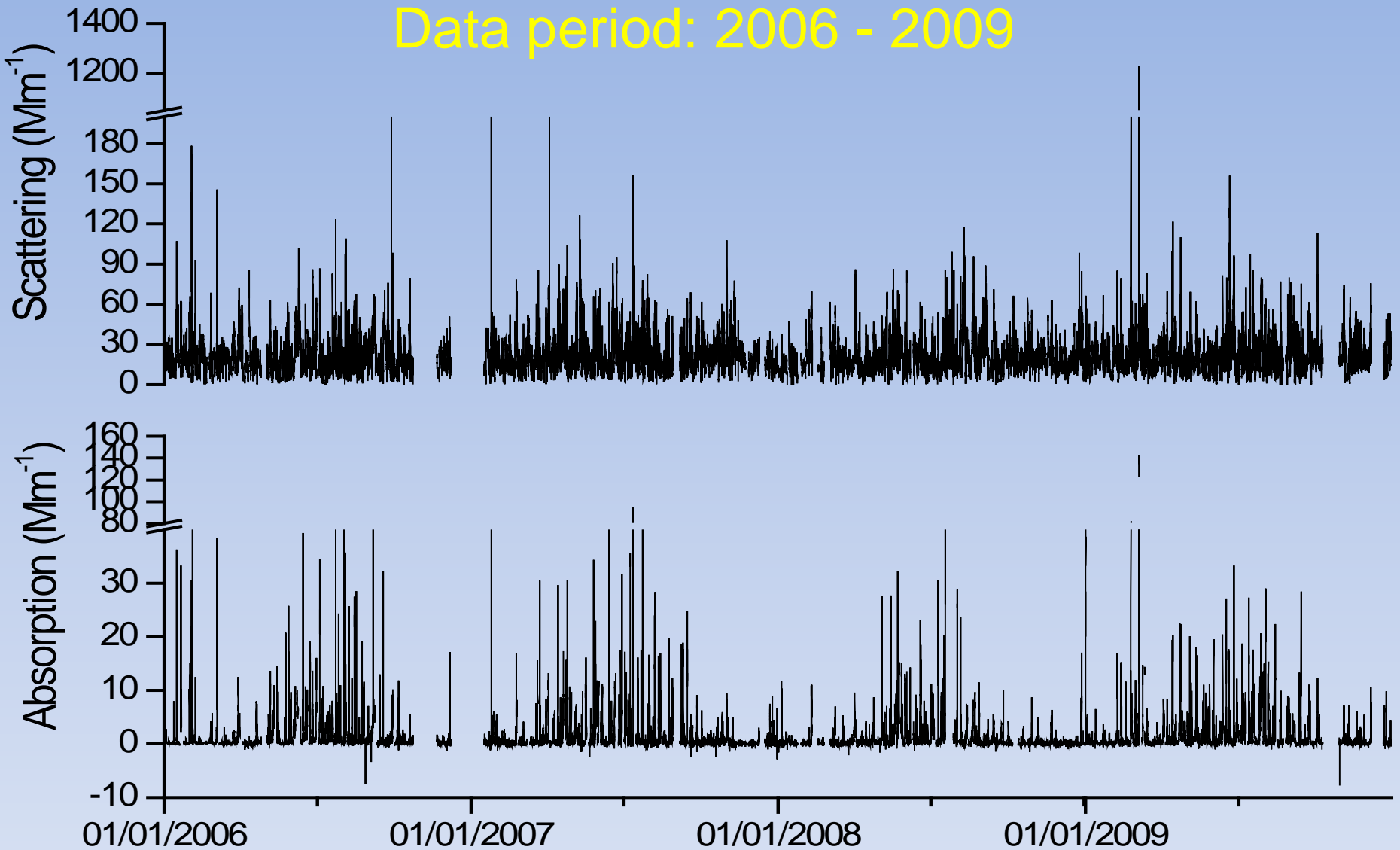
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*NOAA ESRL*



# Aerosol sampling & analysis



Data period: 2006 - 2009

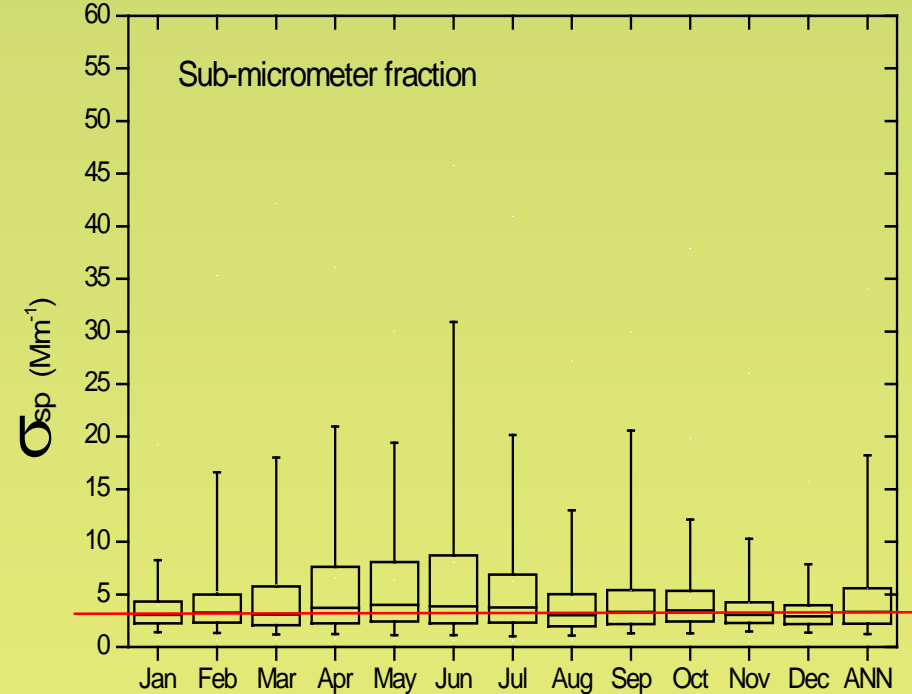
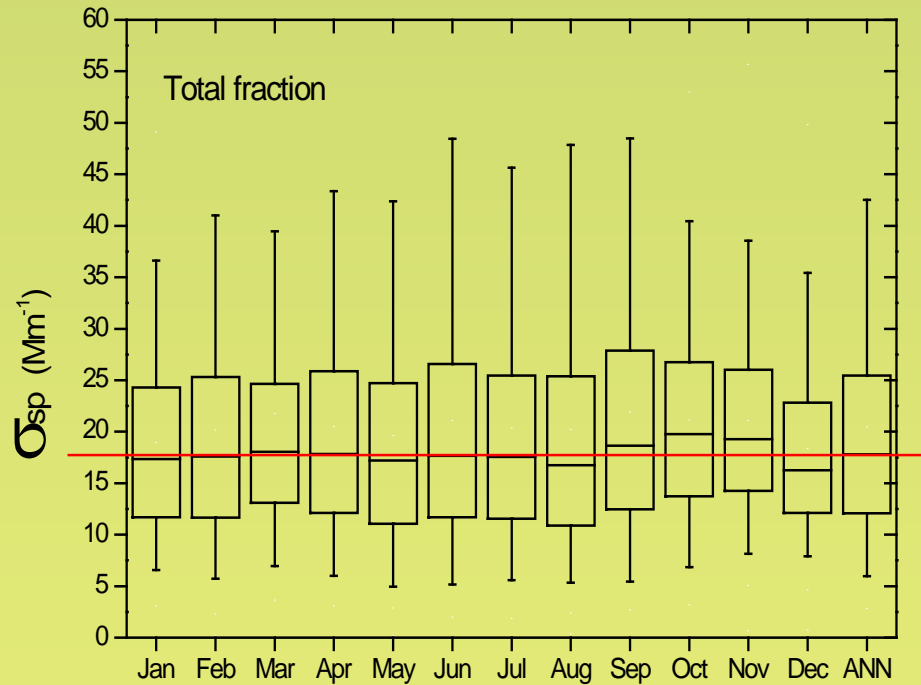


“In the Lord we trust.... all others must provide data!”

anonymous

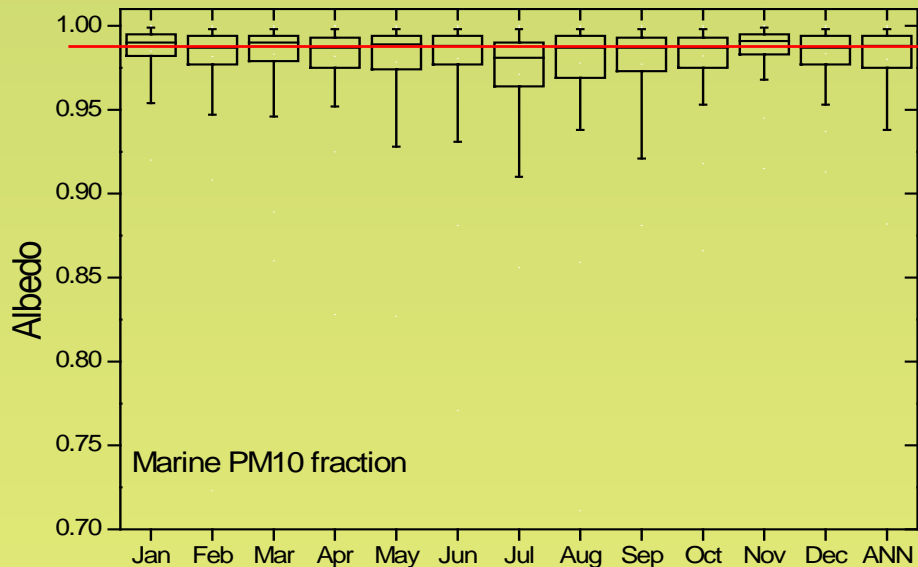
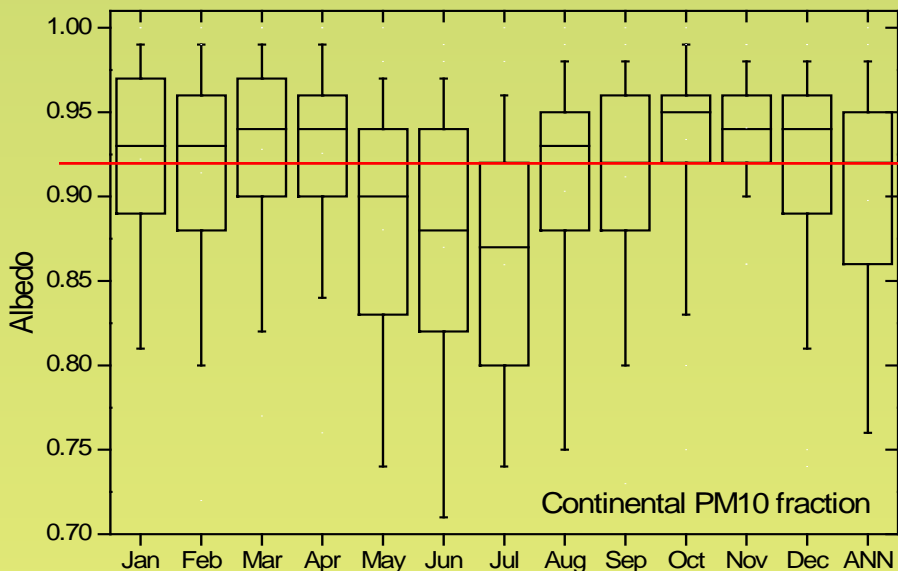


# Aerosol scattering ( $\sigma_{sp}$ )



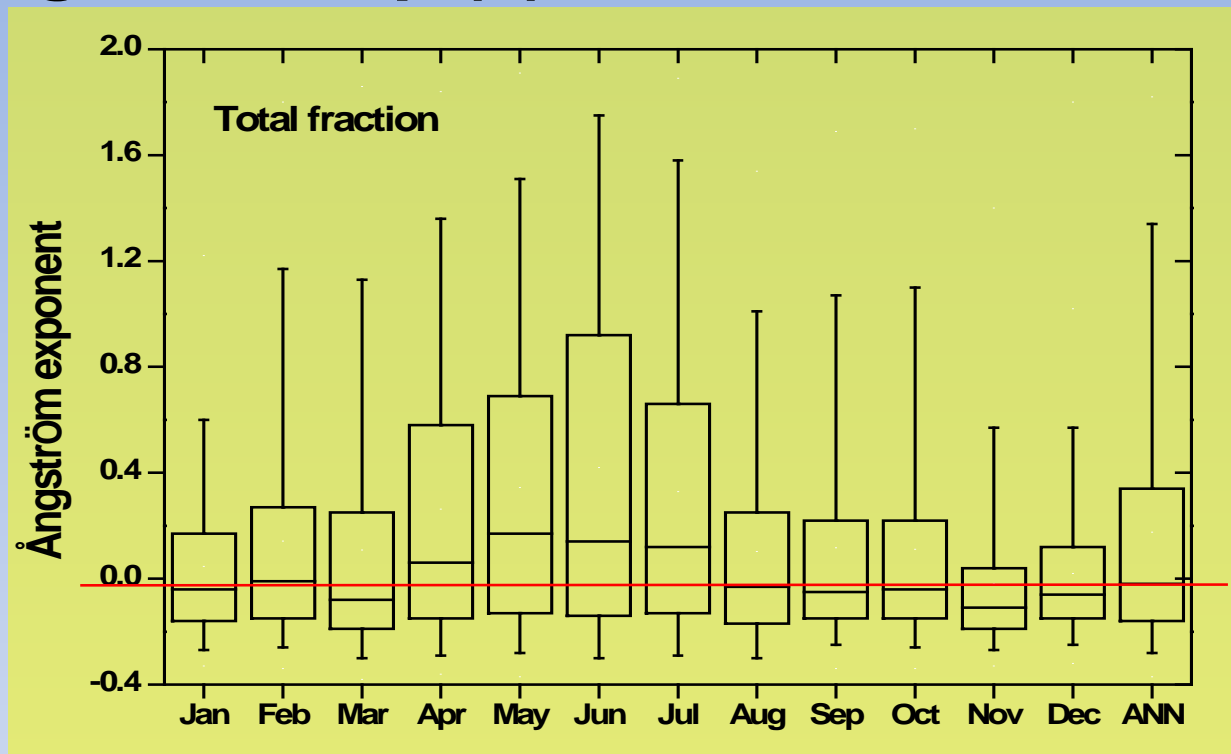
- ❖ Seasonal pattern is discernable in sub- $\mu m$  data, contrasting the total fraction
- ❖ Annual mean aerosol scattering of PM10 fraction is 6x higher than PM1 fraction
- ❖ For the PM1 fraction, Summer months (DJF) have the least variation in aerosol light scattering ( $2 - 4 Mm^{-1}$ ) – highest wind speeds; mainly from south (marine organic particles & predominantly sea salt particles?)
- ❖ Late Autumn - Winter months (MJJA) are dominated by urban particles; also evident in  $\omega$  and  $\dot{a}$

# Single Scattering Albedo ( $\omega_0$ ): Continental vs Marine fraction



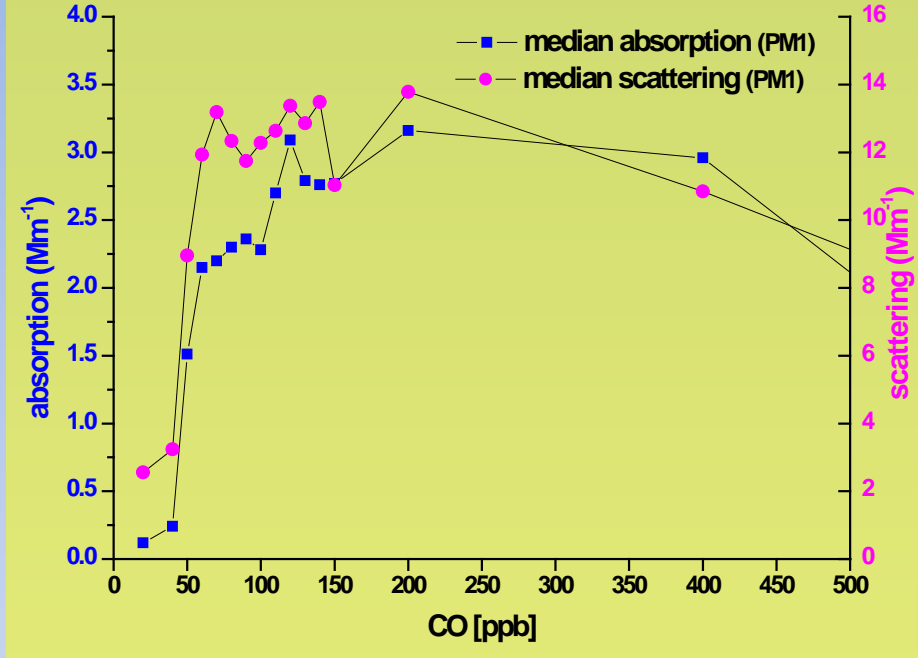
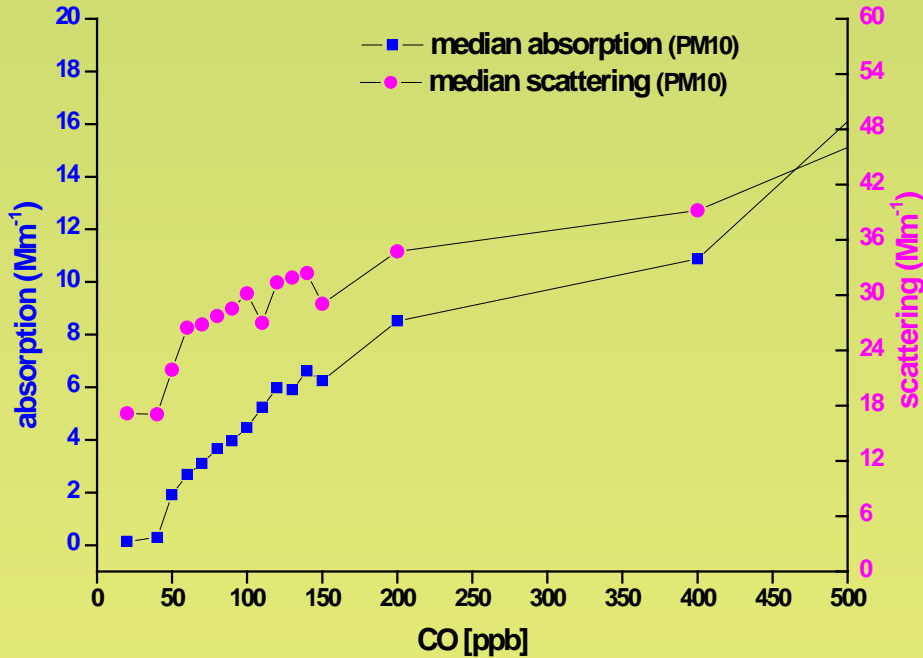
- ❖ Continental =  $^{222}\text{Rn} > 1200 \text{ mBeq/m}^3$  Marine =  $^{222}\text{Rn} < 350 \text{ mBeq/m}^3$
- ❖ Both Continental & Marine albedo fractions display an annual cycle
- ❖ Continental annual mean  $\omega_0 \sim 0.92$  compared to  $\sim 0.99$  for marine
- ❖ Marine fraction – generally, most of seasonal variation is in the 5<sup>th</sup> – 25<sup>th</sup> percentile of data compared to 5<sup>th</sup> – 95<sup>th</sup> percentile of Continental fraction
- ❖ From May – Sep, Continental albedo fraction contains Urban signal & instances of Biomass burning (resulting in higher amounts of absorbing aerosols)
- ❖ Complex to decipher when considering  $\omega_0$  alone – deal with frontal systems as well as incidences of inversion trapping of pollutants

# Ångström exp ( $\text{\AA}$ )<sub>550/700 nm</sub> – PM10 fraction



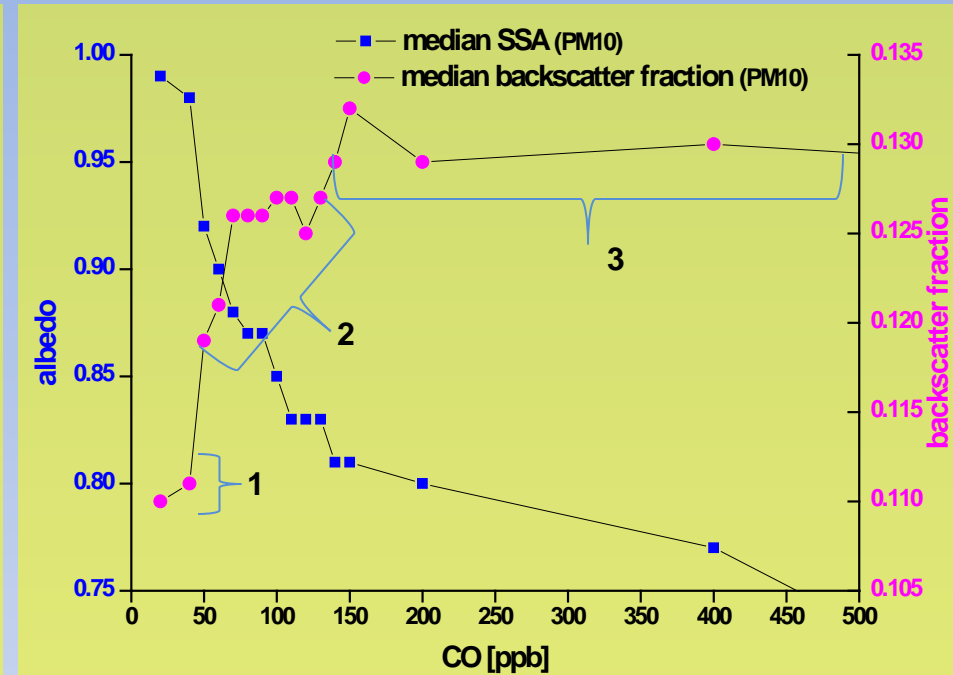
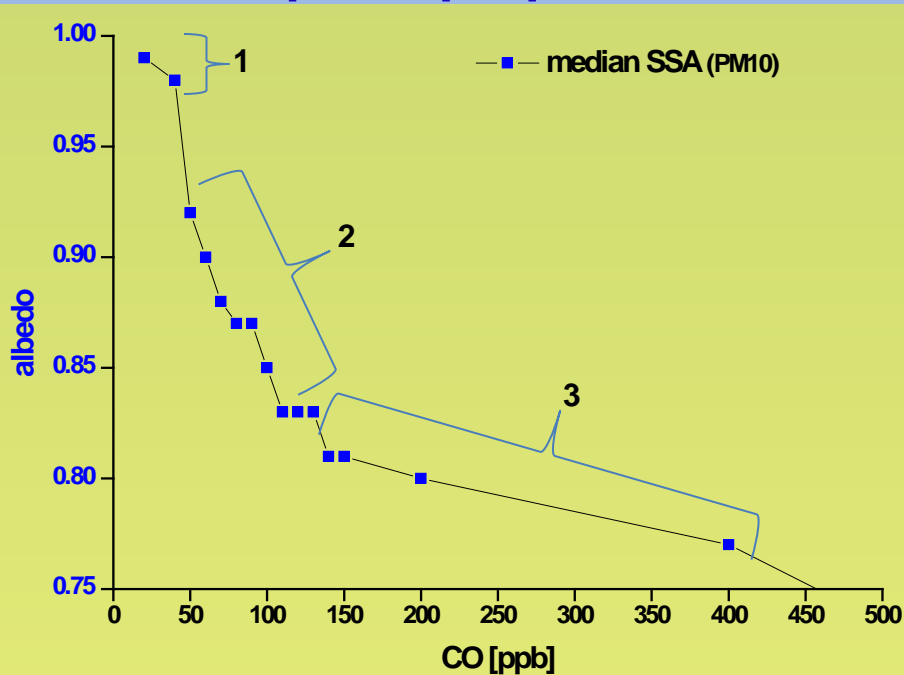
- ❖ Smaller  $\text{\AA}$  indicative of large particles & bigger  $\text{\AA}$  indicate smaller particle sizes
- ❖ Strong evidence of a seasonal pattern in data set
- ❖ Smaller particles dominate in late autumn - winter months (AMJJ when mean  $\text{\AA} > 0$ ) & larger particles in summer season (NDJFM) when mean  $\text{\AA} \leq 0$  (mainly sea-salt dominance)
- ❖ Strongest seasonal pattern observed in the 75<sup>th</sup> - 95<sup>th</sup> percentile data

# Aerosol optical properties as f(x) of Carbon monoxide distribution



- ❖ CO is a good indicator of anthropogenic influence(s): at CPT baseline CO ranges from 40 – 70 ppb
- ❖ Both absorption and scattering follows similar increases with increasing [CO]
- ❖ For the sub- $\mu m$  data – sudden increase in absorption & scattering as CO increase above 70 ppb

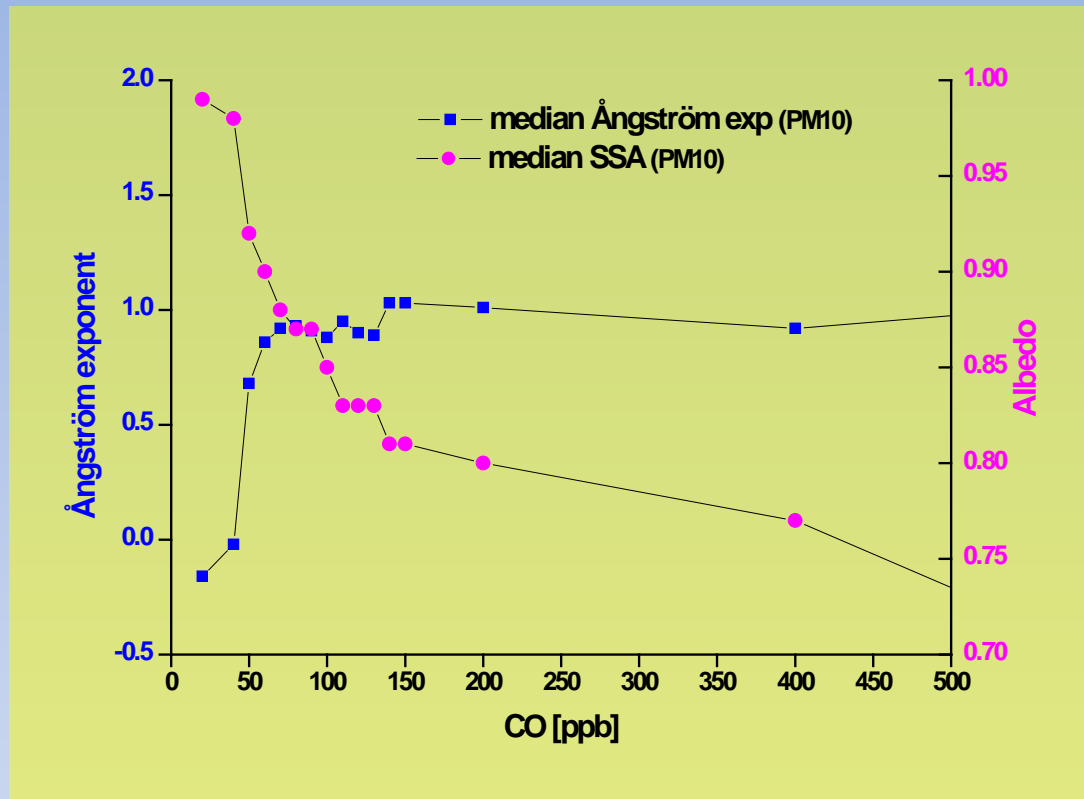
# Aerosol optical properties as f(x) of Carbon monoxide distribution (cont.)



- ❖ Changes in albedo is very sensitive to variations in CO
- ❖ For CO < 50 ppb, albedo corresponds to ~0.97 – 1.0 (particles of mostly of reflective nature)
- ❖ For CO < 130 ppb, albedo ~0.93 – 0.83 (darker particles more absorptive in nature)
- ❖ 3 albedo “ranges” representing (1) clean, marine air (2) mixed urban / continental / biomass (3) urban combustion & biomass burning
- ❖ Backscatter fraction is another indicator of particle size: ratio of  $\sigma_{bsp\_G} / \sigma_{sp\_G}$
- ❖ As CO increases above baseline, smaller particle sizes become more prominent

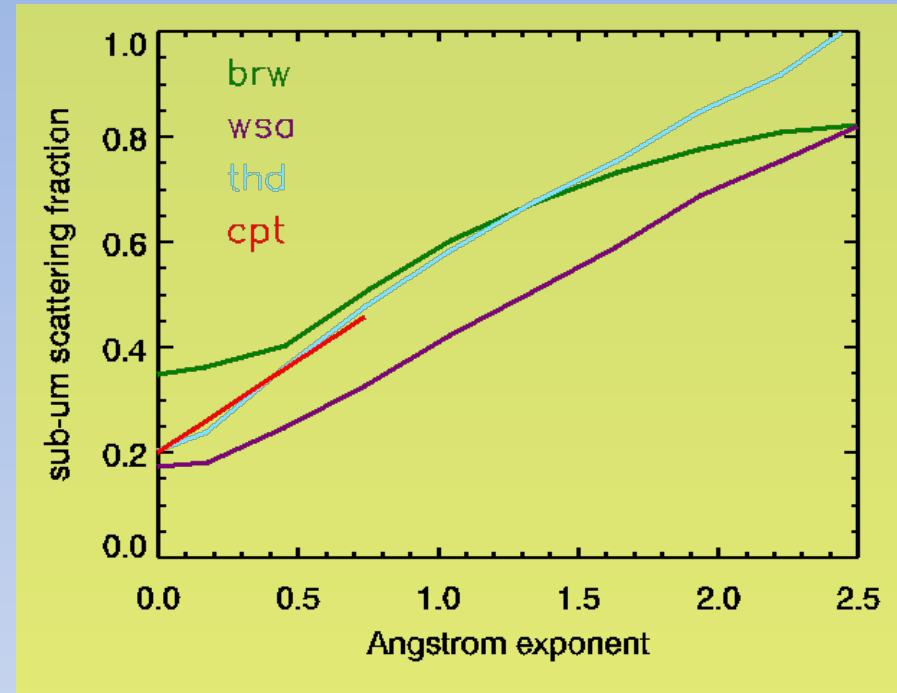
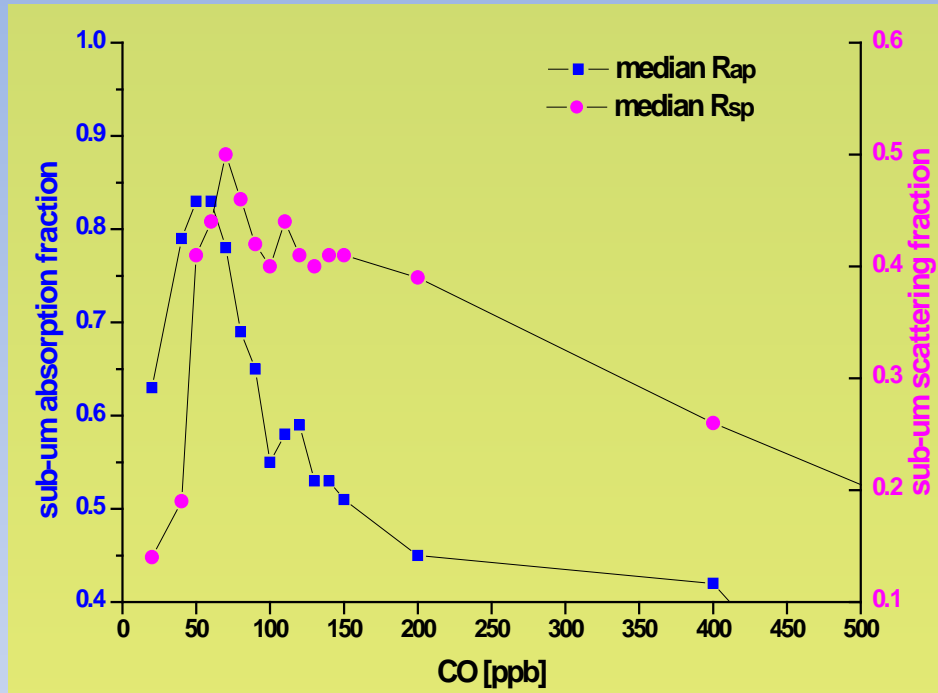


# Aerosol optical properties as f(x) of Carbon monoxide distribution (cont.)



- ❖ Changes in Ångström exp is also very sensitive & abrupt when [CO] changes to polluted conditions
- ❖ For CO > 50 ppb, very clear shift from larger particle sizes to smaller sizes
- ❖ At CPT [CO] < 50 ppb  $\approx \text{Å} \leq 0$  and [CO] > 100 ppb  $\approx \text{Å} = 1$
- ❖ Particle size remain fairly constant on [CO] ~100 - 400 ppb; SSA indicate increase in absorptive nature over the same range
- ❖ For the sub10 $\mu\text{m}$  data, median  $\text{Å} \approx 1$  associated with a SSA of 0.76 – 0.86

# Aerosol optical properties as f(x) of Carbon monoxide distribution (cont.)



- ❖ Sub- $\mu\text{m}$  scattering fraction increases on [CO]  $\sim$ 20 -100 ppb then levels off at 0.4
- ❖ Sub- $\mu\text{m}$  absorption fraction decreases from 0.85 – 0.45 in 70+ ppb [CO] range
- ❖ Comparison of CPT with 3 marine stations (brw=Barrow, Alaska; wsa=Sable island, Nova Scotia; thd=Trinidad head, California).
- ❖ As Rsp increases (more scattering due to sub-um particles) angstrom exponent (10um, red/green wavelength pair) also increases

# Summary & Conclusions

- Observed aerosol climatology largely driven by summer – winter wind regimes
- Large range of albedo ( $\omega_0$ ) values for CPT: e.g. continental winter-time total fraction varies between 0.70 – 0.97
- Ångström exp annual cycle clearly reflects the dominance of large, sea salt aerosols in summer months and smaller, combustion related aerosols in winter
- At [CO] above 100 ppb, small particles play a significant role in defining CPT albedo & Ångström exponents
- Complimentary trace gas data (e.g. CO &  $^{222}\text{Rn}$ ) useful tool in characterizing overall aerosol climatology at Cape Point



**THANK YOU**

**ANY QUESTIONS?**

